In the calibration of the model, the precipitation data from the Redickville station was used for the Mad River. Similarly the Shelburne precipitation station was used for the Boyne River and the Shanty Bay precipitation station for Willow Creek.

As indicated previously, the NWS snowmelt routine requires several input parameters. The most important and most sensitive parameters are:

MFMAX - Maximum non-rain melt factor

MFMIN - Minimum non-rain melt factor

UADJ - Mean wind function value during rain on snow periods

SI - Areal water equivalent above which there is always complete areal snow cover (mm)

In the calibration of the model various ranges of parameter were used. (Ref. 19).

#### These are:

MFMAX 0.004 - 0.009 MFMIN 0.0018 - 0.0035 UADJ 0.017 - 0.057 SI 64 mm - 128 mm

The best overall results were obtained as indicated below:

MFMAX 0.005 MFMIN 0.0018 UADJ 0.057 SI 127 mm

For the Boyne River catchment it was found that the recession constant (K) computed by the Williams equations had to be reduced by 30% to 0.7 time

the equation values. The presence of the reservoir within Earl Rowe Park upstream of the hydrometric station complicated the calibration procedure.

Plots of the observed and simulated hydrographs for each of the spring calibration events are presented in the Figures 3.6(a) to 3.6(f).

#### 3.2.5 Model Validation

#### 3.2.5.1 General

Events used over the summer and the spring period for the validation of the QUALHYMO model were selected from the computer plots of daily flows that were recorded over the 1975 to 1979 period. Hourly recorded discharges for the selected events thereafter formed a basis of comparison with simulated flows.

### 3.2.5.2 Lumped Models

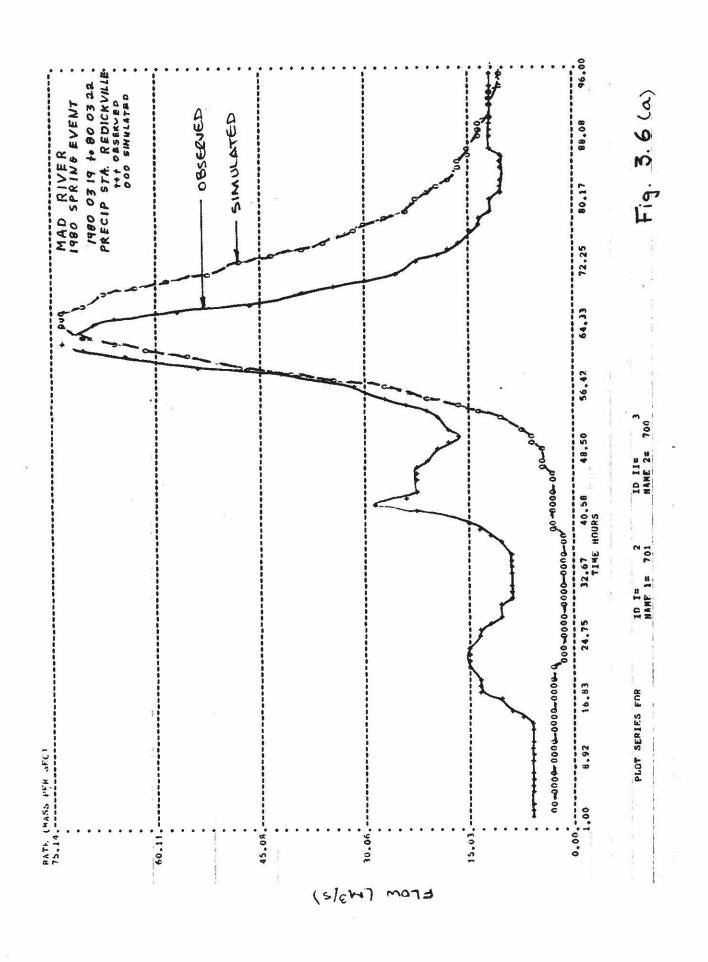
Suitable validation events over the summer period were very sparse between 1975 and 1979. Only one event was found for the Mad River on September 26, 1977 (Table 3.7). Validation for the lumped catchment model for this event proved to be successful (Figure 3.7). A high flow event was also measured on the Boyne River at the Earl Rowe Park gauge; however, hourly streamflow records are not available and validation could not be carried out.

Validation events (Table 3.9) for the spring period were more numerous since annual peak flows usually occur during the freshet as a result of snowmelt, rainfall on saturated ground, or a combination of both. Plots (Figures 3.8 (a) to 3.8(f)) of the observed and simulated hydrographs for each of the spring validation events indicate a close agreement with the exception of the March 25, 1976 event. The observed flow peaks on Willow Creek were checked against those measured at adjoining watersheds and were found not to reflect the general pattern of higher runoff during the initial event.

TABLE 3.9 SPRING VALIDATION EVENTS

	PEAK FLOW (m1/s)				OBSERVED RAINFALL (mm)					OBSERVED SHOW WATER EQUIVALENT (mm)					SIMULATED		
WATERCOURSE	Date	     Observed	Simulated	Type of Event	Date	Barrie	Shanty Bay	Allis- ton	Shel- burne	Redick- ville	li Date	Eden- vale	Colwell	Totten- ham	   Hono   Centre	Maple Valley	Snow Water Equivalent (mm)
Had River at	21 Mar. 1976	102	88.5	Snowmelt	23 Har. 24 Har. 25 Har.	0 2.4 0	0	0 2.4 0	0 0 0	0	15 Feb. 15 Mar. D1 Apr.	71 135 <	76 112 AS	0   0   0   SUME 0	71 112	91 132	133 114 3.4
Glencairn	13 Har. 1977	57.9	81.5	Snowmelt   13 Mar   28	14 Feb. 28 Feb. 15 Mar.	122 99 <	94 84	No Data Trace NO DATA-	83.8 76.0		134 132 25						
Willow Creek above Little Lake	25 Har. 1976	31.1	9.31	Snowmelt	23 Mar. 24 Mar. 25 Mar.	0 2.4 D	0	0 2.4 0	0	0 0 0	15 Feb. 15 Mar.	71 135	76 112	0	71 112	91 132	150 142
	24 Har. 1979	26.5	41.5	Rainfall on Snowmalt	23 Mar. 24 Mar.	4.0 14.0	4.2 13.8	14.8	=	0 16.2	15 Mar. 01 Apr.	130	89	0	31	91	185 14.8
Boyne River at Earl Rowe Park	13 Har. 1971	46.4	42.0	Rainfall on Snowmelt	12 Mar. 13 Mar.	18.1	18.5	20.1	16.1	19.1	14 Feb. 28 Feb. 15 Mar.	122 99 <	94 84	No Data Trace	83.8 76.0	122	87.2 90.7 10
	12 Apr. 1978	25.1	33.6	Rainfall on Snowmalt	10 Mar. 11 Mar.	6.7 10.0	7.8	7.4	10.8 6.5	0 22	15 Mar. D1 Apr. 15 Apr.	191 183 0	155 178 0	91 5 0	130 175 0	150 183 0	156.8 132.4 52.6

<sup>1</sup> Anomaly in Historical Streamflow Data



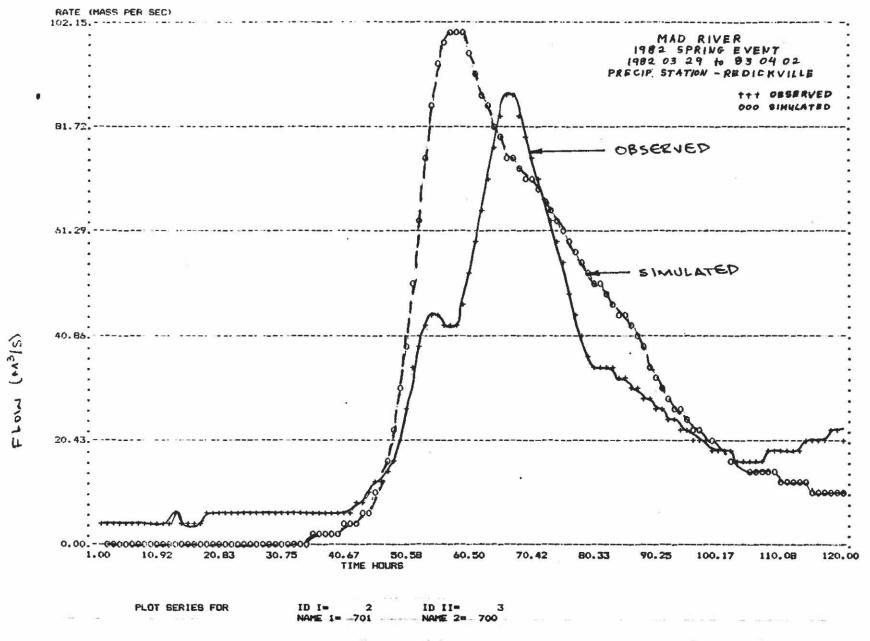
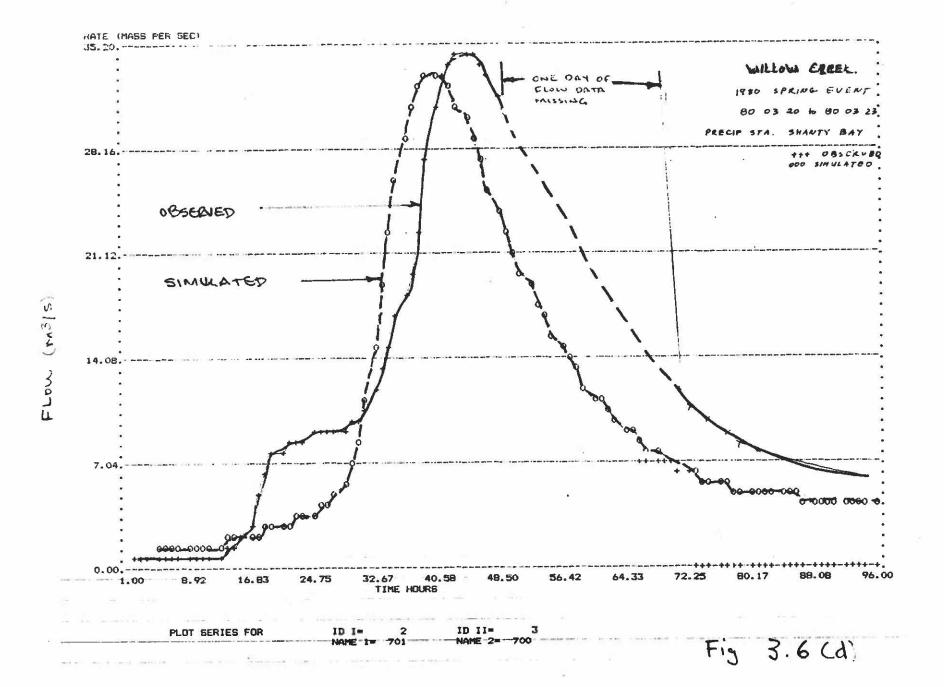
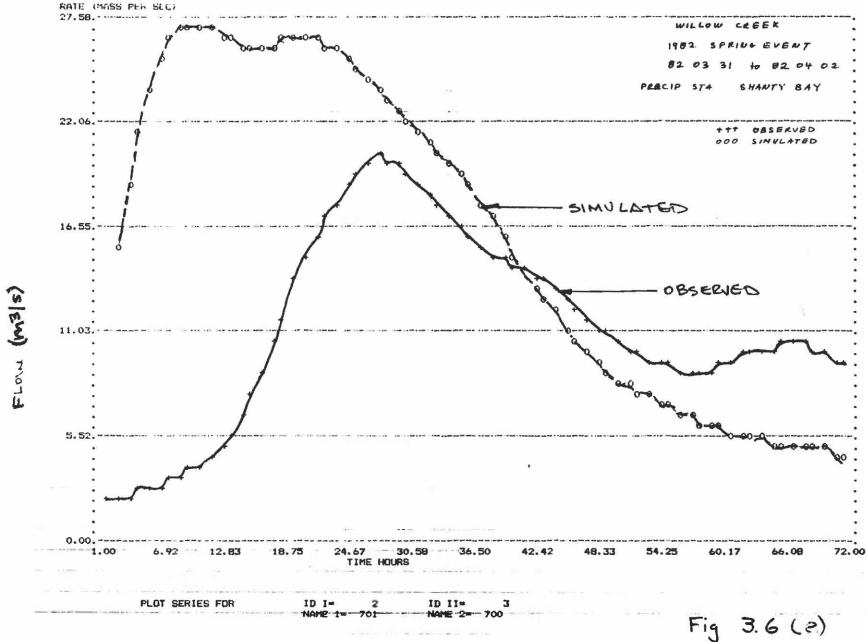


Fig 3.6 (b)

Fig 3.6(c)





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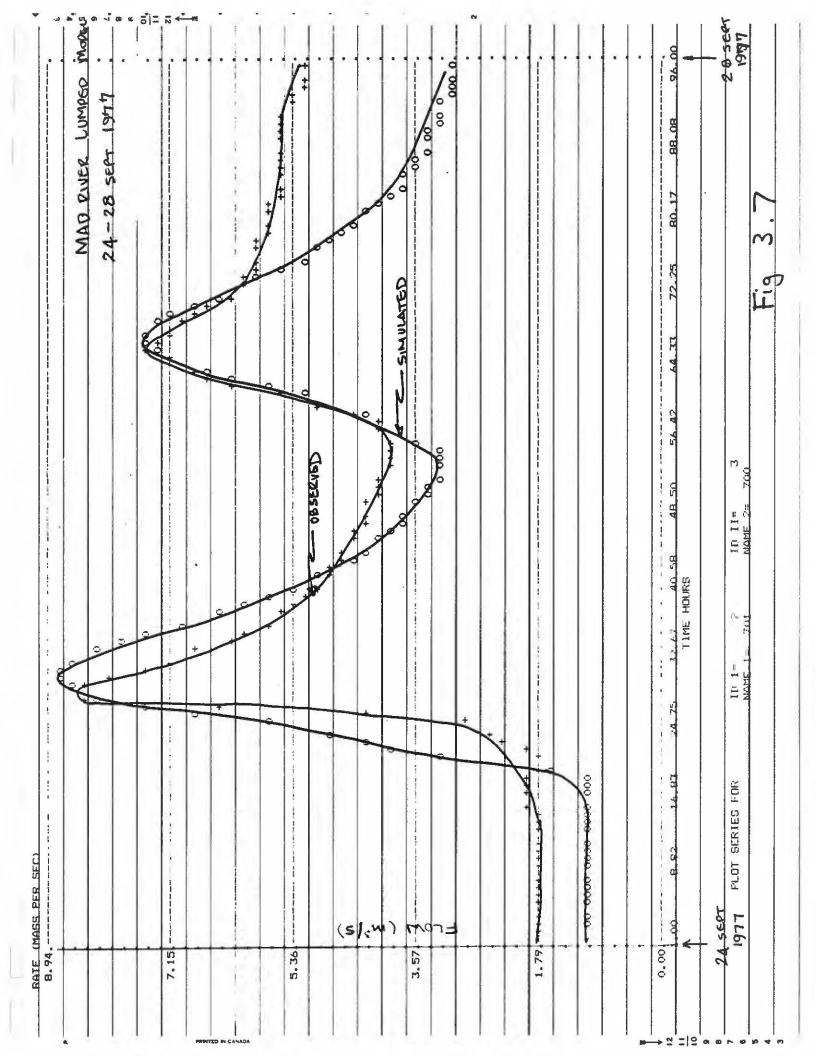


Fig 3.8 (a)

Fig 3.8 (b)

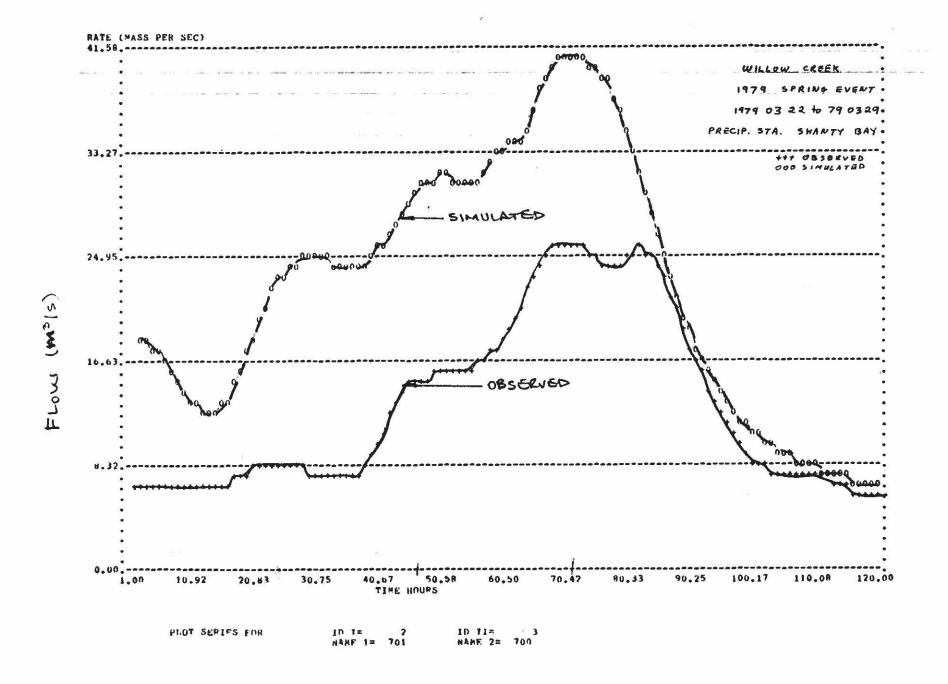
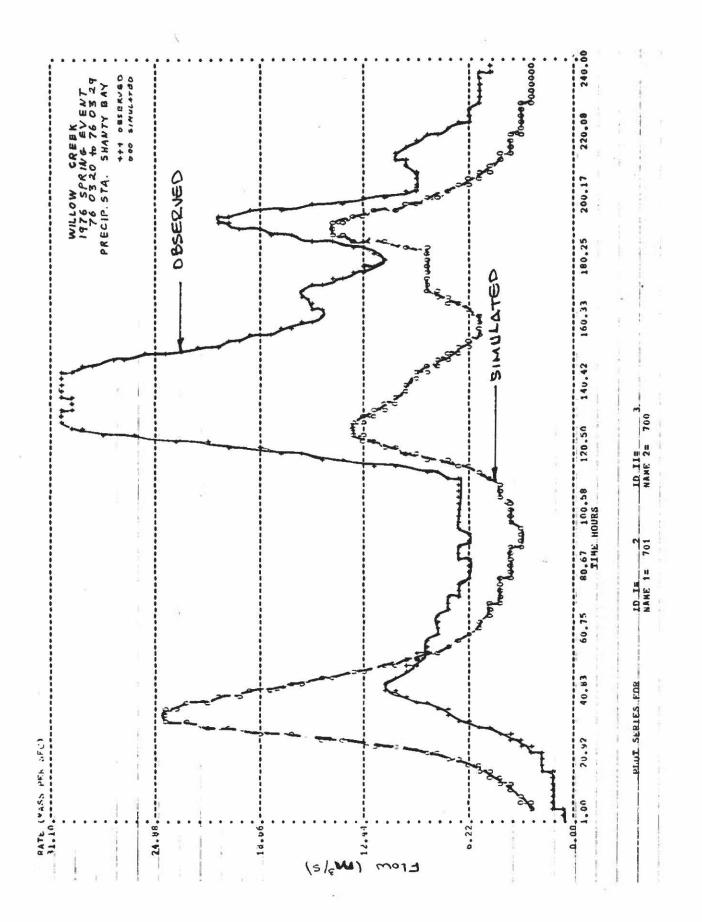


Fig. 3.8 (c)



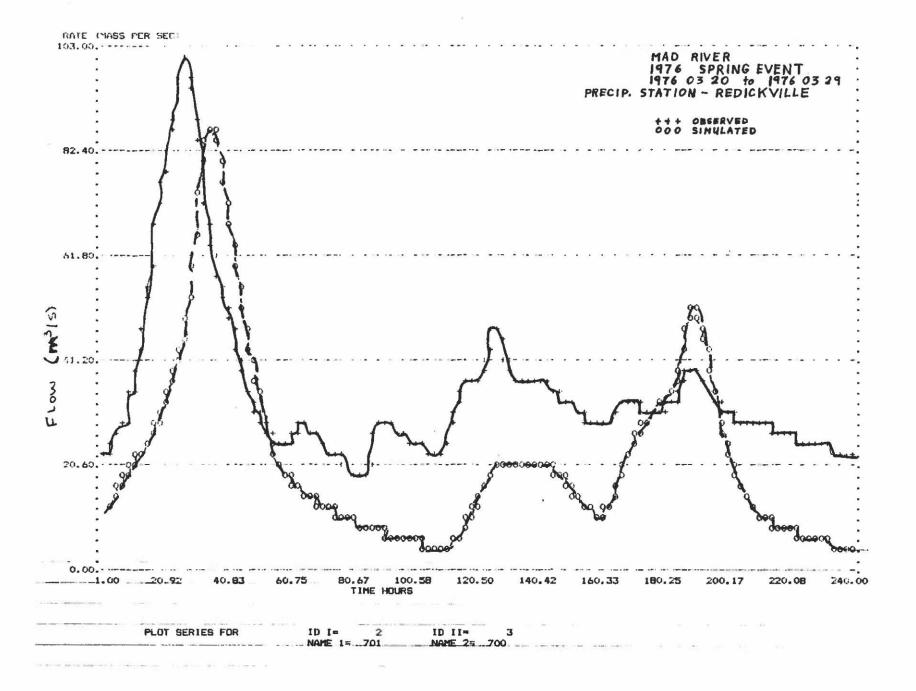


Fig. 3.8(e)

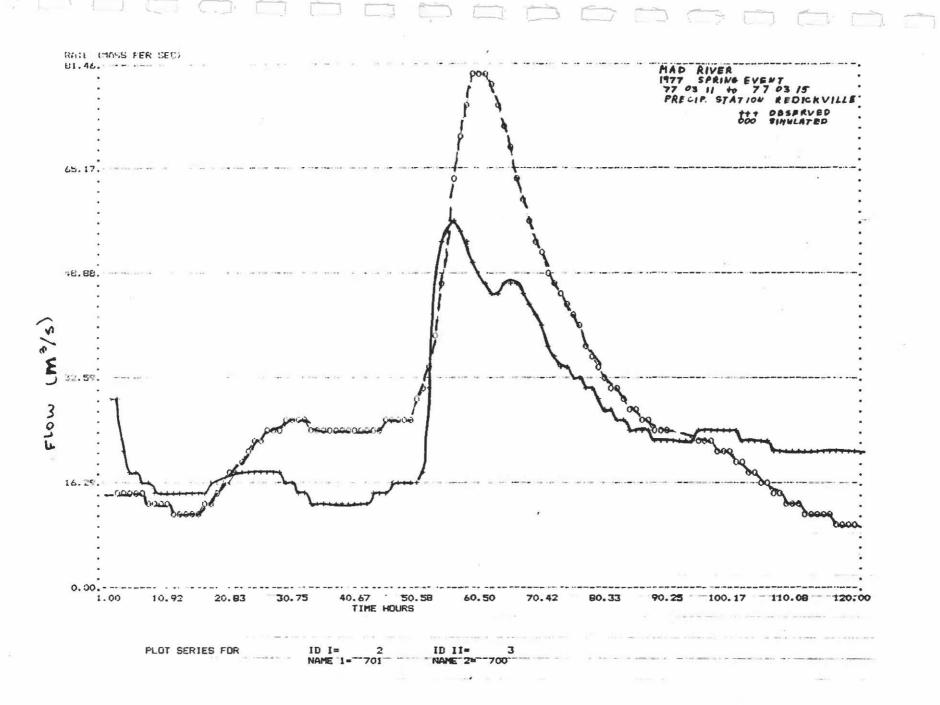


Fig. 3.8(f)

Since discussions with Water Survey of Canada indicated that the Willow Creek gauge above Little Lake is subject to backwater affects from the lake, this event was dismissed as an anomaly.

## 3.2.5.3 Discretized Model

The calibration/validation for the summer and spring events discussed in the previous sections was undertaken using lumped models. In order to confirm the above results, the discretized models for the Mad River, Boyne River, Pine River, Willow Creek and Nottawasaga River catchments were run for specific events.

Hydrologic model schematic for the Nottawasaga River and its tributaries and the Nottawasaga Bay watercourses are presented in Appendix K.

The results using the discretized models for some of the summer events are shown in Figure 3.9(a) to 3.9(c). For the spring and summer events, the comparison in peak flows is summarized in Table 3.10.

Since the lumped and discretized model produce similar results, this comparison provided a firm basis during the subsequent evaluation of frequency based design flows to use the lumped models for the simulation of the 22 years of historical flows and the discretized model to distribute the flow to tributary sub-catchments.

A further analysis of the sensitivity of flows with the time step used in the Variable Storage Coefficient routing was carried out (Table 3.10a) for the discretized model. Little variation in peak flow was noted for a shorter routing interval (15 minutes) than used in the model calibration and validation (one hour); therefore, the one hour time step was considered acceptable for application of the discretized model.

### 3.2.5.4 Historical Flood Peaks

As further validation of the QUALHYMO models, the annual peak flows were simulated for the 1963 to 1984 period and compared with observed discharges in a scatter diagram indicating individual events and by frequency analyses.

TABLE 3.10

Comparison of Results Using Lumped and
Discretized QUALHYMO Models

Simulated Peak Flow Observed Peak Discretized Lumped Model Mode 1 Catchment Flow Date  $(m^3/s)$  $(m^3/s)$  $(m^3/s)$ 74.4 75.0 78.2 21 March Mad River near Glencairn 1980 51.0 46.4 42.0 Boyne River 13 March at Earl Rowe Park 1977 33.5 38.6 Willow Creek 21 March 35.3 above Little Lake 1980 5.0 3.9 4.4 Beeton Creek 29 July near Tottenham 1980 15.7 19.2 Pine River near 29 July 21.7 Everett 1980 66.7 66.3 96.6 Nottawasaga River 29 July near Baxter 1980

<sup>(1)</sup> Observed baseflow of 5.0  $m^3/s$  was added to simulated peak flow of 14.2  $m^3/s$ 

## TABLE 3.10a

# SENSITIVITY TESTING OF ROUTING EFFECT USING 0.25 HOURS AND 1.0 HOUR TIME STEP FOR TIMMINS STORM (All flows in m³/s)

		0.25 HOU	R TIME STEP	1.0 HOUR TIME STEP			
REACH NO.	FLOW POINT	PEAK FLOW	ROUTED PEAK FLOW	PEAK FLOW	ROUTED PEAK FLOW		
	1070	1657.5	-	1648.0	-		
44	1072	1682.8	1646.2	1674.3	1639.9		
45	1074	1743.8	1682.5	1733.2	1672.7		
46	1078	1844.6	1741.4	1827.2	1727.6		
48	480	1867.9	1832.1	1851.5	1819.1		
49		1007.5	1844.9	1001.0	1833.2		

